Relation among the processes of erosion-sedimentation-pollution in soils from the Distrito Pecuario “Alturas de Nazareno”, Cuba

J. M. Febles González1, N. Amaral Sobrinho2, Y. Pérez López3, J.H. Zoffoli2, M.O. Lima Magalhães2 and N. Guedes2

1Universidad de La Habana, Zapata y G, Vedado, Plaza de la Revolución, Ciudad de La Habana, C.P.: 10 400.
2Universidade Federal Rural do Rio de Janeiro. Br 465 km 7, Seropédica - R J, Brasil
3Universidad Agraria de La Habana “Fructuoso Rodríguez Pérez” Carretera Jamaica-Tapaste, km 21/2, San José de Las Lajas, Cuba
Email: febles@rect.uh.cu

The geological environments and their influence on the processes of erosion-sedimentation-pollution in soils due to heavy metals in the Distrito Pecuario “Alturas de Nazareno” were evaluated, with a functional biological diversity supported by the natural regeneration of the tree component. The work of pedological prospection: identified four types of soils, described and sampled at 0-20 cm and 20-40 cm deep. It was confirmed the prevalence of erosive processes and the possible access of the grass roots into a layer of peculiar alteration (level 25 - 30 cm) in “Alturas de Nazareno”, with considerable proportions of NiO2, Zn, Cd, Cr and Fe. In the Polje of San José, there was a preeminence of the sedimentation-pollution from the surface (level 0 - 30 cm), associated to toxic organic compounds, nutrients and organic matter, where any not controlled action in the cattle systems could move the heavy metals present in the soil to the pastures and the water provided to the animals.

Key words: heavy metals, trace elements, geo-chemical bank, reference values, polje

The heavy metals (HM) are naturally present within the soils. The increase of their concentrations can occur due to natural processes or anthropogenic activities (Guilherme et al. 2005 and Amaral Sobrinho et al. 2009).

The total natural concentration of HM within the soils depends mainly on the original material, formation processes and the proportion of the components of the solid phase. This concentration is less evident in soils formed over sediments (Fadigas et al. 2006 and Utermann et al. 2006).

Researches carried out in the carcasic regions of Havana province, in the sub-basin Mampostón and in the Distrito Pecuario “Alturas de Nazareno”, have revealed the soil losses due to the erosion in ferrallitic environments, which propitiate the sequential degradation of the most productive soils of Cuba (Febles et al. 2012).

In this context, there are still unspecific regularities and genetic bonds of the processes of sedimentation and contamination with the manifestations of soil erosion, whether they have geological character in less disturbed environments or induced by the agricultural and livestock activity (Febles et al. 2010). As it is known, the HM can be present naturally within the soils, even though they have not had anthropic disturbance (Guilherme et al. 2005). Starting from this point, this study has the objective of making an integral evaluation of the geological environments and their influence on the processes of erosion-sedimentation-pollution in soils due to heavy metals in the Distrito Pecuario “Alturas de Nazareno”

Materials and Methods

The Distrito Pecuario “Alturas de Nazareno” is located in the center of Mayabeque province, in the Western part of Cuba. It has an extension of 71261.5 ha. It starts with a small tectonic-structured height, like horst type, with flat top (Iturralde-Vinent 2011), pertaining to the regions of Bejucal-Madruga-Coliseo Heights. They descend abruptly through a well defined escarpment into a denudative plain, developed over a rock bed (of plinth), slightly dissected (H = 100 – 120 m).

It receives, approximately, between 76 and 80 % of the annual precipitations of the province (INSMET 2012), which are superior to 1,580 mm per year. This fact makes this agro-ecosystem to be considered as one of the most humid in Cuba. The minimum absolute temperatures per year in the territory show average values of 19.5 ºC. The maximum absolute temperatures per year are of 26.33 ºC and the annual mean temperature is 23.47 ºC.

The main soils (table 1) are characterized by a functional biological diversity, supported by the natural regeneration of the tree component, with “healing” species like marabou (Dichrotechys glomerata), aroma (Acacia farnesiana) and palm trees (Roystonea regia), which act as an induced fallow land and favor certain resilience or stability in the processes, with some independence of the energy value of the relief. Nevertheless, during the 80’s, the cover was treated, in some sectors, with a non differentiated management technology.

Twelve main profiles were characterized. Soil samples were taken from the superior, medium and inferior flexures of the micro-relief, in order to examine the dynamic, manifestation and intensity of the erosion
through a comparative-descriptive analysis, at depths of 0-20 cm and 20-40 cm.

The samples were pounded, grinded up to 210 µm particle size. They were subjected to an extract of aqua regia (3HNO₃ + 1HCl) to quantify the pseudo-total contents (do not destroy the structures of the silicates) (ISO 1995). The fractions of metals in the soil samples were determined by the “European Communities Bureau of Reference” (BCR), according to Sahuquillo et al. (1999).

The geochemical fractions were: F1 extracted with a solution of acetic acid 0.11 mol L⁻¹ (soluble acid fraction), F2 extracted with a solution of hydroxylamine hydrochloride 0.1 mol L⁻¹ (fraction linked to the Fe and Mn oxides), F3 extracted with the solutions of hydrogen peroxide 8.8 mol L⁻¹ and ammonium acetate 0.11 mol L⁻¹ (soluble acid fraction), and F4 extracted with aqua regia – Σ from the previous fractions (residual fraction). The analyses of HM were performed with five analytic repetitions per sample. The mean values of the repetitions per each type of soil were considered.

The concentrations of the elements in the extracts resulting from the soil were quantified with an inductively coupled plasma optic emission spectrometer (ICP-OES), Pekín Elmer trademark, ÓPTIMA 3 000 model, with LD¹ of (mg kg⁻¹) of 0.508 and LQ (mg kg⁻¹) of 1.69 for Zn; LD (mg kg⁻¹) of 0.026 and LQ (mg kg⁻¹) of 0.097 for Cd; LD (mg kg⁻¹) of 0.20 and LQ (mg kg⁻¹) of 0.067 for Pb; LD (mg kg⁻¹) of 0.01 and LQ (mg kg⁻¹) of 0.05 for Fe; LD (mg kg⁻¹) of 0.01 and LQ (mg kg⁻¹) of 0.020 for Mn; LD (mg kg⁻¹) of 0.03. The certified material of reference NIST SRM 2709a (San Joaquin Soil, with Zn concentrations of 103 ± 28 mg kg⁻¹; Cd, 0.371 ± 0.002 mg kg⁻¹, Pb 17.3 ± 0.1; Mn of 529 ± 18 mg kg⁻¹; Fe of 33.6 ± 0.7 g kg⁻¹) was used for validating the determination of the pseudo-total content of Zn, Cd, Fe, Mn, Pb and Cr in the soil.

**Results and Discussion**

Due to that the edaphic cover “stores” features and properties inherited from past geological and climatic phases, which are not in balance with the current edaphogenic processes (Guerasimov 1983), this analysis started from the paleo-geographic evolution of sub-district of the Habana-Matanzas Heights, which has been very complex. It has also been necessary to reflect on the possible origin of the HM currently present in the soils of the Distrito Pecuario “Alturas de Nazareno”. The paleo-geographic evolution of this sub-district started from the sediments of the Laramide basement, which bended profoundly with frequent thrust faults during the development stage of the isles bow (Iturralde 1996).

This sub-district can be considered as part of the Continental Paleomargin of Bahamas, where appeared the carbonated rocks from the inferior Eocene (P₂) to the inferior Miocene (N₁) (Albear and Iturralde - Vinent 1985), in a complex combination with other very different rocks (table 2), like tufa, marlstones and sandstones (Campos et al. 2004). There are also, underlying, the ophiolites from the Capdevila (P₂) formation.

In these enclaves (Febles et al. 2009), the balance of morphogenesis-pedogenesis depend on the morpho-structural, and very specially, on the chemical-mineralogical composition of rocks. This process can be very slow, as happens with the ophiolites, in which the morphogenesis can even inhibit the edaphogenesis on the rocks, due to the erosion of the eluvia while they are forming. It can also be fast, like in the tufa stones, in which it is possible to go from the denudation to the formation of deep soils.

There are geological and geomorphologic peculiarities in these environments, which influence directly on the erosion processes, whether they are geological or induced by the agricultural and livestock activity, with more extended actions through time regarding other environments. A significant characteristic of the soils from the Distrito Pecuario “Alturas de Nazareno” is, precisely, their moving towards the bedrock. There is no C horizon, and it goes suddenly from B horizon to the underlying limestone (figure 1). This happens in most cases through a very thin layer, determined by its reaction to acid, which is similar to the peculiar alteration described by Lamoureux (1972) in Lebanon and by Bosch et al. (1984) in ferrallitic soils from Cuba.

Specifically, this layer contains, at around

### Table 1. Area occupied by the main types of soil in the Distrito Pecuario “Alturas de Nazareno”, Mayabeque province

<table>
<thead>
<tr>
<th></th>
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</thead>
<tbody>
<tr>
<td>Typical red ferrallitic</td>
<td>38 106.496</td>
<td>Ferralsol Rhodic</td>
<td>Ferrallitic grouping. Red type</td>
</tr>
<tr>
<td>Carbonated brown</td>
<td>24 658.760</td>
<td>Cambisol Calcic</td>
<td>Brown sialitic grouping. Brown type</td>
</tr>
<tr>
<td>Red rendzine</td>
<td>5 823.987</td>
<td>Feozem - Calcric-Rhodic - Skeletic</td>
<td>Humic grouping. Rendzine type</td>
</tr>
<tr>
<td>Brown red Fersialitic</td>
<td>2 634.142</td>
<td>Cambisol Chromic-Humic</td>
<td>Fersialitic grouping. Brown red</td>
</tr>
</tbody>
</table>

¹LD: Limit of detection and LQ: Limit of quantification, of the analytic methods used.
Table 2. Main rocks in the formation of ferrallitic eluvia in soils from the Distrito Pecuario “Alturas de Nazareno”, Mayabeque province

<table>
<thead>
<tr>
<th>Rocks</th>
<th>Characteristics</th>
<th>References</th>
</tr>
</thead>
<tbody>
<tr>
<td>Aleurites</td>
<td>Constituted by angular and sub-angular grains of quartz, with some fragments of plagioclases, mica and dark minerals. The relative proportion of the different lyotologies is variable, with a predominance of limes and the limes-aleuritas-argilitas rythm.</td>
<td>Albear and Iturralde (1985)</td>
</tr>
<tr>
<td>Ophiolites</td>
<td>Capdevila formation (subjacent to Nazareno group) Inferior Eocene.</td>
<td>Bronnimann and Rigassi (1963)</td>
</tr>
<tr>
<td>Greywackes</td>
<td>They are associations of ultramafic, mafic and mafic-volcanic rocks. Constituted by mica, feldspar, and other grain constituents, including quartz (although in lower proportions). All those elements are united by an also dentritic cement.</td>
<td>Bronnimann and Rigassi (1963)</td>
</tr>
<tr>
<td>Relictus-organic and crystalline limes</td>
<td>They appear at open sky quarries in Camoa and Somorrostro, with a light grey color and black.</td>
<td>Bronnimann y Rigassi (1963)</td>
</tr>
</tbody>
</table>

20-30 cm deep in the non carbonated residues, considerable proportions of NiO₂, Zn, Cd, Cr, Fe, among other elements. According to Camacho et al. (1980) and Camacho et al. (1985), for some limestone studied in Cuba, these elements enrich the cortexes of metheorization with HM, where the greatest proportion of forages roots and is concentrated. This fact has been confirmed in the researched region, because there is a good correspondence between the mineralogy of rocks and soils. They are much related to the degree of impurities of limestone, their speed of dissolution and the time passed, which has been confirmed by Ortega (1984) and Jaimez and Ortega (2001) in similar territories.

The previously mentioned has been also confirmed by the comparative analysis of the natural values of the HM in the main soils from the “Alturas de Nazareno” (table 3), regarding the reference values, used internationally (Lista holandesa 2000, Kabata-Pendias and Pendias 2001, CETESB 2005 and Fadigas et al. 2006). There is also evidence that the brown red fersialitic, typical red and carbonated brown ferrallitic soils present natural values over those limits.

In some cases, these soils can be classified as “contaminated”. However, these concentrations are found naturally, due to the presence of these elements in the minerals that form the rocks and the geological environments of formation, which coincides with those found by Fadigas et al. (2002) in studies carried out in Brazil.

The geo-chemical behavior of the HM in these soils shows some similarities for the different fractions (table 4). The Cu, Zn, Ni, Fe and Cr showed higher percentages in the residual fraction (F4), which indicates this elements are found in the most stable geo-chemical fraction, with the least solubility and bioavailability, generally related to primary and secondary minerals, which can contain metals within their crystalline structure, hardly released in less disturbed environments or in biostasia (Dang et al. 2002).

The geo-chemical distribution of Mn is mostly found in the fraction associated with the oxides of Fe and Mn (F2). On this matter, Rodriguez et al. (2009) point out that the association of metals with oxides does not guarantee the immobilization in the surface because these elements are still unstable in this fraction and they may be released by the reduction of oxides and, consequently, cause a negative effect on the soil and on the biota (Cholpecka 1996).
Table 4. Sequential extraction of pseudo-total HM in the main types of soils at 0-40 cm deep. Distrito Pecuario “Alturas de Nazareno”, Mayabeque province

<table>
<thead>
<tr>
<th>Fraction</th>
<th>Cu (mg kg⁻¹)</th>
<th>Zn (mg kg⁻¹)</th>
<th>Mn (mg kg⁻¹)</th>
<th>Ni (mg kg⁻¹)</th>
<th>Cd (mg kg⁻¹)</th>
<th>Pb (mg kg⁻¹)</th>
<th>Fe (mg kg⁻¹)</th>
<th>Cr (mg kg⁻¹)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Brown red fersialitic</td>
<td></td>
<td></td>
<td></td>
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<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>F1</td>
<td>1.8 (2)</td>
<td>0.0 (0)</td>
<td>56.9 (2)</td>
<td>2.5 (2)</td>
<td>0.9 (30)</td>
<td>9.2 (18)</td>
<td>1.8 (0.1)</td>
<td>0.0 (0)</td>
</tr>
<tr>
<td>F2</td>
<td>11.0 (13)</td>
<td>7.5 (4)</td>
<td>1771.1 (68)</td>
<td>22.3 (14)</td>
<td>0.7 (24)</td>
<td>20.3 (40)</td>
<td>611.8 (2)</td>
<td>0.6 (1)</td>
</tr>
<tr>
<td>F3</td>
<td>12.9 (15)</td>
<td>3.7 (2)</td>
<td>361.5 (14)</td>
<td>16.2 (10)</td>
<td>0.8 (26)</td>
<td>36.2 (72)</td>
<td>58.2 (0.1)</td>
<td>11.3 (13)</td>
</tr>
<tr>
<td>F4</td>
<td>58.5 (70)</td>
<td>160.3 (93)</td>
<td>409.7 (16)</td>
<td>120.1 (75)</td>
<td>0.6 (19)</td>
<td>0.3 (1)</td>
<td>27964.7 (97)</td>
<td>50.7 (56)</td>
</tr>
<tr>
<td>Typical red ferrallitic</td>
<td></td>
<td></td>
<td></td>
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<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>F1</td>
<td>0.8 (1)</td>
<td>8.6 (2)</td>
<td>349.9 (13)</td>
<td>2.0 (1)</td>
<td>2.2 (19)</td>
<td>2.8 (4)</td>
<td>0.0 (0)</td>
<td>0.0 (0)</td>
</tr>
<tr>
<td>F2</td>
<td>1.0 (1)</td>
<td>14.8 (3)</td>
<td>1722.7 (63)</td>
<td>6.9 (2)</td>
<td>4.3 (38)</td>
<td>0.5 (1)</td>
<td>914.9 (2)</td>
<td>0.1 (0)</td>
</tr>
<tr>
<td>F3</td>
<td>14.5 (11)</td>
<td>13.8 (3)</td>
<td>98.8 (4)</td>
<td>8.7 (4)</td>
<td>1.4 (13)</td>
<td>0.0 (0)</td>
<td>576.1 (1)</td>
<td>12.3 (4)</td>
</tr>
<tr>
<td>F4</td>
<td>113.2 (87)</td>
<td>431.8 (92)</td>
<td>554.0 (20)</td>
<td>290.5 (93)</td>
<td>3.3 (29)</td>
<td>62.0 (95)</td>
<td>46629.7 (97)</td>
<td>277.9 (96)</td>
</tr>
<tr>
<td>Carbonated brown</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>F1</td>
<td>1.7 (4)</td>
<td>2.9 (4)</td>
<td>149.7 (15)</td>
<td>4.4 (8)</td>
<td>0.9 (40)</td>
<td>12.3 (28)</td>
<td>3.4 (0)</td>
<td>0.7 (3)</td>
</tr>
<tr>
<td>F2</td>
<td>0.5 (1)</td>
<td>1.7 (3)</td>
<td>572.8 (58)</td>
<td>2.8 (5)</td>
<td>0.6 (24)</td>
<td>4.8 (11)</td>
<td>0.0 (0)</td>
<td>0.0 (0)</td>
</tr>
<tr>
<td>F3</td>
<td>6.1 (14)</td>
<td>8.4 (13)</td>
<td>214.2 (22)</td>
<td>12.0 (21)</td>
<td>0.4 (16)</td>
<td>12.8 (29)</td>
<td>664.1 (3)</td>
<td>5.5 (16)</td>
</tr>
<tr>
<td>F4</td>
<td>36.1 (81)</td>
<td>52.8 (80)</td>
<td>55.3 (6)</td>
<td>38.4 (67)</td>
<td>0.5 (20)</td>
<td>14.4 (32)</td>
<td>20899.1 (97)</td>
<td>27.0 (81)</td>
</tr>
</tbody>
</table>

Values in parenthesis represent the percentage regarding the total; F1: soluble acid fraction; F2: fraction linked to oxides of Fe and Mn; F3: fraction linked to organic matter; F4: residual fraction.

Table 3. Natural values of HM in soils from the Distrito Pecuario “Alturas de Nazareno”, with relation to the reference values proposed by the international literature

<table>
<thead>
<tr>
<th>Soil Type</th>
<th>Cu (mg kg⁻¹)</th>
<th>Zn (mg kg⁻¹)</th>
<th>Mn (mg kg⁻¹)</th>
<th>Ni (mg kg⁻¹)</th>
<th>Cd (mg kg⁻¹)</th>
<th>Pb (mg kg⁻¹)</th>
<th>Fe (mg kg⁻¹)</th>
<th>Cr (mg kg⁻¹)</th>
</tr>
</thead>
<tbody>
<tr>
<td>B.R.F.¹</td>
<td>84.0</td>
<td>171.5</td>
<td>2599.0</td>
<td>161.1</td>
<td>3.1</td>
<td>50.3</td>
<td>28639.0</td>
<td>90.3</td>
</tr>
<tr>
<td>T.R.F.²</td>
<td>129.5</td>
<td>469.1</td>
<td>2725.0</td>
<td>308.1</td>
<td>11.3</td>
<td>65.4</td>
<td>48120.0</td>
<td>290.0</td>
</tr>
<tr>
<td>B.C.³</td>
<td>44.4</td>
<td>65.9</td>
<td>992.0</td>
<td>57.7</td>
<td>2.4</td>
<td>44.4</td>
<td>21567.0</td>
<td>33.2</td>
</tr>
<tr>
<td>Q.R.⁴</td>
<td>35.0</td>
<td>60.0</td>
<td>*</td>
<td>13.0</td>
<td>0.5</td>
<td>17.0</td>
<td>*</td>
<td>*</td>
</tr>
<tr>
<td>Prev.⁵</td>
<td>60.0</td>
<td>300.0</td>
<td>*</td>
<td>30.0</td>
<td>1.3</td>
<td>72.0</td>
<td>*</td>
<td>*</td>
</tr>
<tr>
<td>Res.⁶</td>
<td>200.0</td>
<td>450.0</td>
<td>*</td>
<td>70.0</td>
<td>3.0</td>
<td>180.0</td>
<td>*</td>
<td>*</td>
</tr>
<tr>
<td>B.C. ³</td>
<td>2.0 - 119</td>
<td>6.0 - 79</td>
<td>*</td>
<td>5.0 - 35</td>
<td>0.3 - 1.5</td>
<td>3.0 - 40</td>
<td>*</td>
<td>19.0 - 65</td>
</tr>
<tr>
<td>Fadigas⁸</td>
<td>0.3 - 495</td>
<td>1.5 - 264</td>
<td>80 - 1315</td>
<td>0.7 - 269</td>
<td>0.005 - 2.4</td>
<td>0.5 - 135</td>
<td>*</td>
<td>6.0 - 80</td>
</tr>
<tr>
<td>Holland²</td>
<td>36.0</td>
<td>140.0</td>
<td>*</td>
<td>35.0</td>
<td>0.8</td>
<td>85.0</td>
<td>*</td>
<td>100.0</td>
</tr>
</tbody>
</table>


The Cd and Pb are distributed in a differentiated manner, according to the type of soil. The brown red fersialitic soil, like the carbonated brown soil, presented the highest tenors of Cd in the acid soluble fraction (F1). That fraction contains the elements of the solution, weakly adsorbed by electrostatic interaction, and precipitated with carbonates (Marín et al. 1997 and Filgueiras et al. 2004). Their high content could be a problem from the environmental point of view, mainly in the fractions of higher lability (Sastre et al. 2001 and He et al. 2005).

Pb presented a higher tenor in the fraction linked to the organic matter (F3) in the brown red fersialitic soil, while for the typical red ferrallitic soil, the highest percent was found on the residual fraction (F4). For the carbonated brown soil, there was no predominance in a determined fraction. Very proximate values were found, and in all the fractions.

In a general evaluation, the majority of the elements in all the soils are mainly distributed in the residual fraction, as a consequence of the low solubility and strong adsorption in the surface of the soil particles (Chaney 1991), and the erosion is the main process for the removal of these elements. This coincides with the informed by Núñez et al. (2006).

“Alturas de Nazareno” is the main source of porous sediments, which sequentially descend towards sectors of the relief of minor hypsometry.
and geo-morphological stability (figure 2), where the competition for the courses of water is decreasing, which provokes the consequent abandonment of sediments (Febles 2007).

Nevertheless, it has been confirmed that, with certain independence of the energy value of the relief, in those sectors where the anthropogenic activity has not taken place, there is a tendency to pedogenesis, with a very slow, positive or practically zero areal removal of the soil fractions (Febles et al. 2007 and Febles et al. 2008).

It is important to say that during the 80’s, the traditional system of soil preparation towards the slopes, the burning and over-grazing increased the acting of the natural processes and induced the development of a group of erosive-gravitational and karstic-erosive processes, which limited the productive capacity and fertility of soils.

This material, while descending, produces the obstruction of the drain lines, as well as the progressive sedimentation of agricultural areas, water bodies and grazing areas, where the grass roots concentrated on the first 30 cm extract the HM, consumed later by the cattle (Febles and Ruiz 2007 and Febles et al. 2012a).

Additionally, the remains of the superficial flow moves towards the polje (figure 3), towards the central zone of San José de Las Lajas municipality, where the heterogeneity of the superficial karstic forms, like dolines, uvales, sump and some other, exert an strict control over the drain (Vega and Febles 2006), with an almost total predominance of sedimentation.

These deposits usually contain high concentrations of HM from even the surface (level 0-30 cm), toxic organic compounds, nutrients and organic matter. It is necessary to have an especial care because any uncontrolled action could activate the toxic compounds retained on the sediments towards urban, aquiferous or cattle areas, which coincides with the reported by Guzmán et al. (2010) in studies carried out in this locality.
It was confirmed the existent relations among the processes of erosion-sedimentation-contamination in different environments and levels of acting in soils from the district “Alturas de Nazareno”. There is, among them, a prevalence of erosive processes and possible access of the grass roots into the layer of peculiar alteration (level 25 - 30 cm), with considerable proportions of Ni, Zn, Cd, Cr and Fe, among other elements. In the polje of San José, it was confirmed the predominance of the sedimentation-pollution, with high concentrations of heavy metals from the surface (level 0 - 30 cm), associated to toxic organic compounds, nutrients and organic matter.

In every soil, the Cu, Cd, Zn, Ni, Fe and Cr elements are mainly distributed in the residual fraction (F4), while in the carbonated brown and brown red fersialitic soils, the Cd and Pb elements are distributed in the soluble acid fraction (F1). This aspect could be a problem from the environmental point of view, mainly in the fractions of higher lability.

The results of this study represent a relevant approach to design more balanced management systems in risk scenarios for the cattle establishment, with a prevalence of grasses and forage. These results can also be useful for any productive purpose in Cuba.

References
Bosch, D., Camacho, E., & Ortega Sastríques, F. 1984. Influencia de las rocas calizas en la génesis de los suelos Ferralíticos Rojos de las llanuras cársicas de Cuba. Ciencia de la Tierra y el Espacio 9: 125
Febles, G. & Ruiz, T.E. 2007. Algunos resultados alcanzados en el desarrollo de los pastos y forrajes en Cuba y elementos de su desempeño en el sector productivo agropecuario. IV Foro Latinoamericano de pastos y forrajes. Instituto de Ciencia Animal, Cuba
Guzmán, A.R. 2010.Caracterización agroecológica y procesos morfogenéticos actuales en áreas agrícolas contaminadas
con metales pesados en el municipio de San José de las Lajas. III Taller Nacional de Biorremediación
INSMET 2012. Datos climáticos. Centro de clima. Instituto de Meteorología. La Habana, Cuba
Núñez, J. E.V., Amaral Sobrinho, N.M.B. & Mazur, N. 2006. Sistema de preparo do solo e acúmulo de metais pesados no solo e na cultura do pimentão (Capsicum annum L.). Ciência Rural. Santa Maria. 36: 113
World Reference Base. 2006. Mapa Mundial de Suelos, escala 1: 30 000 000. World Soil Res.

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